



Geophysik III (Handbuch der Physik Encyclopedia of Physics)

G.M. Nikol'skij, K. Rawer, P. Stubbe, L. Thomas, T. Yonezawa

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83 input to the ions is balanced by a cooling to atomic oxygen, the major neutral constituent, and measurements of the difference between ion and neutral temperature permit the determination of atomic oxygen concentrations. Using this approach, ALCA YDE et al. 6 have shown from data taken above Saint Santin, France, that the atomic oxygen concentration at 200 km is slightly larger in winter than summer. () The molecular concentrations at heights near 200 km can also be derived from a determination of the ratio of the molecular-ion concentration to electron concentration, p (Cox and EVANS 7). It can be shown from the steady state form of the continuity equation for O^+ ions that (20.2) where $Q(O^+)$ represents the photoionization coefficient of oxygen atoms, and k and k_{19} represent the rate coefficients of reactions between O^+ ions and $1S$ molecular nitrogen and oxygen (reactions (12.18) and (12.19)). Then the ratio of atomic oxygen to weighted molecular sum is given by: $(1-p)N_e / [O] = (20.3) Q(O^+) / [k_{19}N_e]$ where N_e represents the electron concentration.

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